



Tidal Rhythmites Depositional System of the Cretaceous Yolde Formation of the Gongola Sub- basin Northern Benue Trough N. E. Nigeria

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Authors' contributions

All the authors participated in field work where data sets used in producing this research article were generated. All authors read and approved the final manuscript.

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ABSTRACT

Facies evaluation carried out on the Yolde Formation at Gabukka locality in the Gongola Sub-basin of the Northern Benue Trough was aimed at enacting its paleo-depositional environment. The formation at this locality revealed an intercalated succession of massive bedded sandstone facies (Sm), planar crossbedded sandstone facies (Sp), ripple laminated sandstone facies (Sr), parallel laminated sandstone facies (Sl) and mudstone facies (Fm). These successions developed a thinning and thickening stratigraphic profile depicting periodic rhythmic signatures indicating deposits of tidal rhythmites. Thickening rhythmic packages are reflective of a spring tide whereas thinning phases are indicative of neap tide. These depositional sequences are genetic to intense tides conditions, thus an indexing a tide dominated oceanographic within the course of development of the Cretaceous Yolde Formation.

Keywords: *Tidal rhythmites; depositional system; Yolde Formation; Gongola Sub-basin; Benue trough.*

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1. INTRODUCTION

Tidal rhythmicity is typically conditioned to tide dominated oceanographic systems and the mid-Cretaceous marine transgression that inundated the West and Central African Rift System set the template for their evolution. The Gongola Sub-basin forming part of this rift system is the north-south trending arm of the Northern Benue Trough that represent the tip of Benue Trough (Fig. 1), evolving primarily due to of separation of the African plate from the South American plate, occurring during the late Jurassic to early cretaceous times. Current, the evolution is this remained controversial, however the rift and pull-apart theories were considered to be more resounding. The rift model theory was proposed at inception by earlier workers and supported to date [1,2,3,4] indicating initiation through tensional regimes induced by mantle plume

convection activities [5,6]. This is opposed to the pull-apart model because of the absences of boundary fault that are proxy to rifting, therefore considered the trough as of strike-slip tectonic origin, as it falls in tune and orientation to the major transcurrent fault systems of the Romanche, Chain and Charcot suture zones [7,8,9]. The opening of the trough is followed by transgressive and regressive phases in the Aptian-Albian times with the Northern Benue Trough characterized by continental depositional regimes. Transgressional activity reached this part of the trough in the Cenomanian, depositing transitional-marine sequences of the Yolde Formation. This researched is aims to evaluate the facies and facies association of this formation at Gabukka stream that represents one of its major outcrops in the Gongola Sub-basin in order to establish depositional model that characterizes its development this locality.

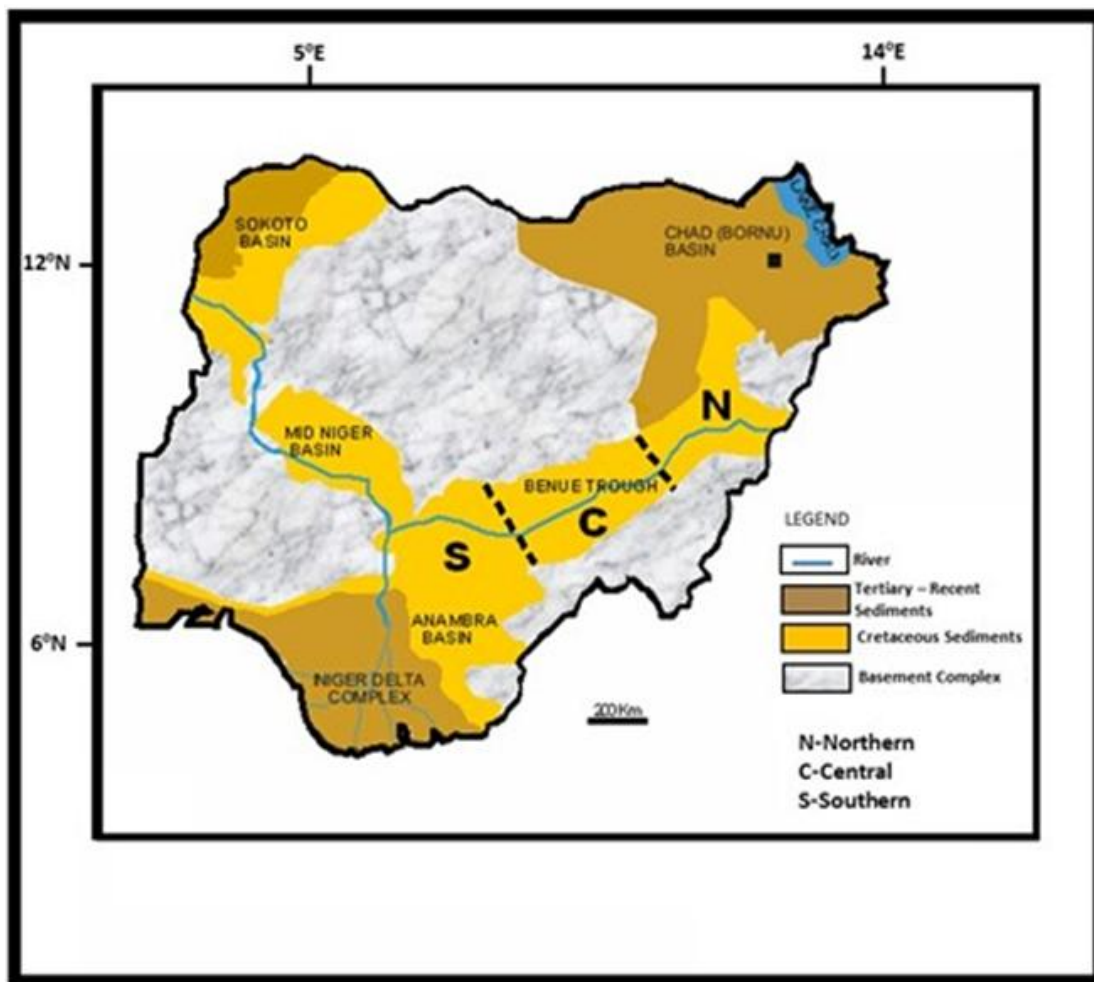


Fig. 1a. Geological map of Nigeria showing the benue trough

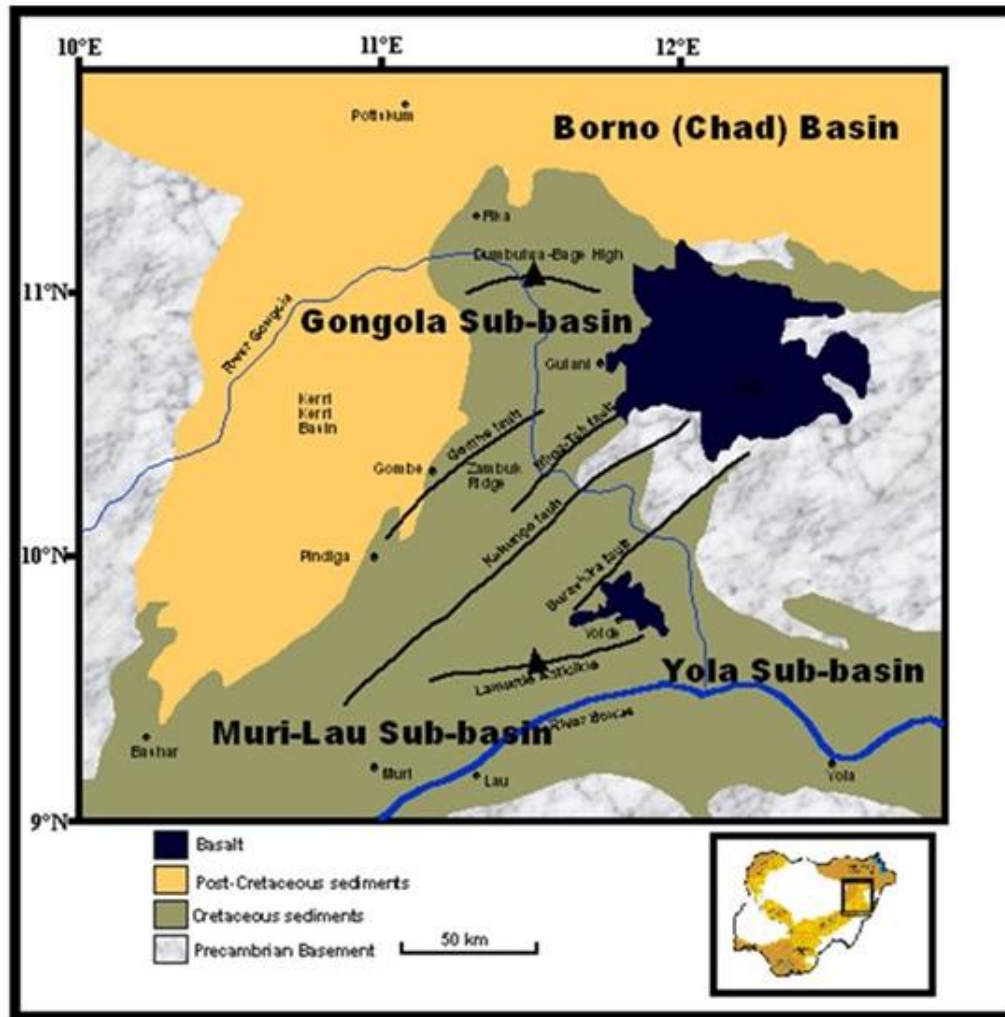


Fig. 1b. Geological map of the Northern Benue Trough

1.1 Geological and Stratigraphic Setting

The Benue Trough of Nigeria is a rift basin in the Central West Africa that extends NNE-SSW for about 1000 km in length and 50-150 km in width [3,10]. The Benue Trough is geographically subdivided into Northern, Central and Southern Benue Trough (Fig. 1). The Northern Benue Trough constitutes of three arms: The N-S striking Gongola Arm, E-W striking Yola Arm and the NE-SW striking Muri-Lau Arm [11]. The Trough contains over 6000 m deep consisting of Cretaceous to Tertiary sediments of which those predating the mid-Santonian have been tectonically deformed, to form major faults and fold systems across the basin. The Bima Group of the Aptian-Albian represents the oldest sedimentary units in the Gongola Sub-basin,

conformably overlying the Basement Complex Rocks (Fig. 2) [12,13,14,15]. The deposition of syn-rift sequences thereof is largely controlled by the horst and graben systems and is represented by the alluvial fan-lacustrine deposits of the Bima I Formation, the lowermost in the group, which is unconformably superposed by the post-rift braided river sequences of the Bima II and III Formations [13,14,15]. This is conformably superposed by the Yolde Formation in the Cenomanian, marking a transitional-marine deposits [16]. This representing the onset of the mid-Cretaceous global marine transgression in the basin [17] and reached its acme in the Turonian depositing the shallow marine shale and limestone sequences of the Kanawa Member of the Pindiga Formation [13,18]. Regressive Sandy Members of the Dumbulwa,

Deba-Fulani and Gulani sandstones conformably followed in the mid-Turonian with decelerating transgressive conditions (Fig. 2) [13,10]. Renewed rising relative sea levels in the late Turonian transcending into the Coniacian and early Santonian set in the deposits of the deep marine blue-black shales of the Fika Member, representing the youngest units of the Pindiga Formation [13,19]. This marine transgression is accompanied by compressional tectonics in the mid-Santonian [20], which resulted from changing orientation of the displacement vectors between the African plate and European/Tethys

plates [21]. This event led to thrusting of the pre-Maastrichtian sediments towards the west of the Gongola Sub-basin, creating an accommodation for the deposition of the Campano-Maastrichtian regressive deltaic sequences of the Gombe Formation [22,19]. The mid-Maastrichtian is also characterized by another phase of compressional event and thereafter followed by the unconformably deposits of the Paleogene fluviolacustrine Kerri Kerri Formation [23,24] (Fig. 2). The Paleogene-Neogene is notable for volcanics, emplaced along the eastern margin of the Gongola Sub-basin [25].

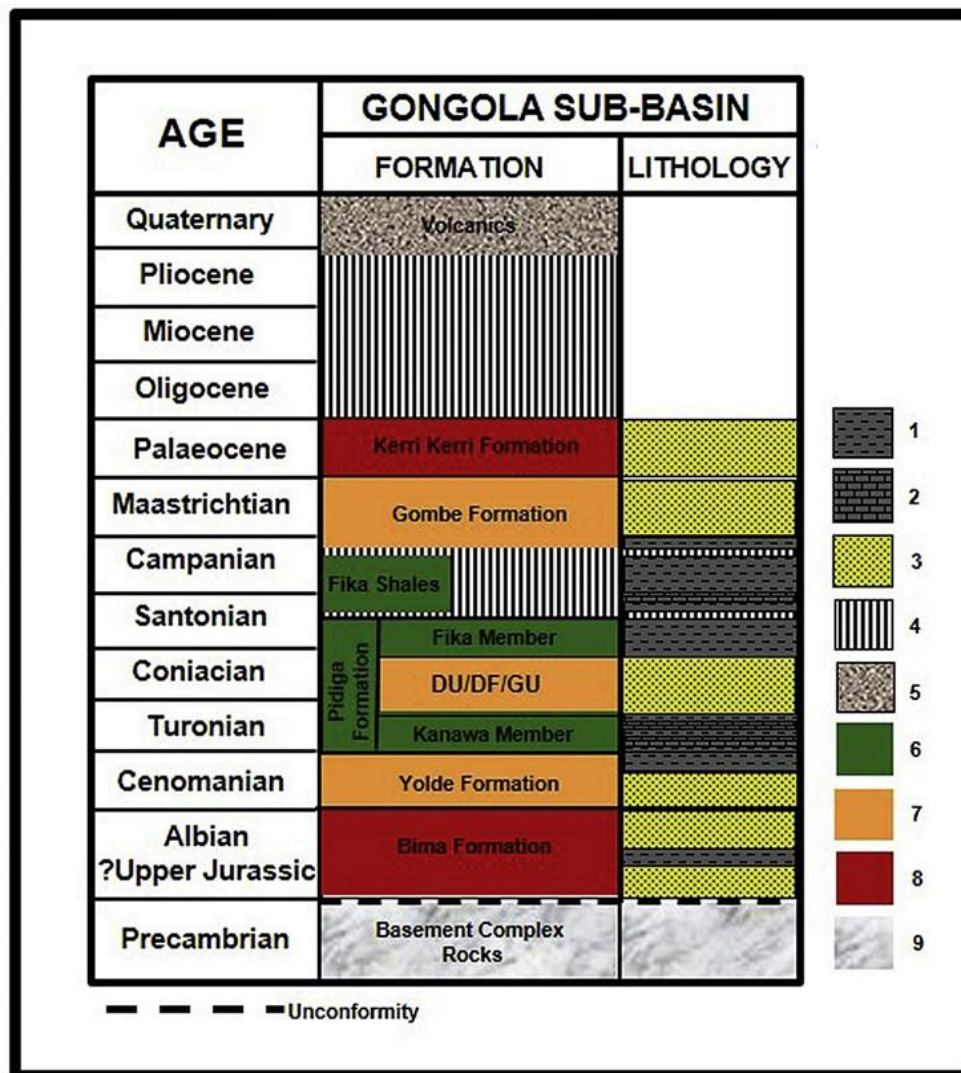


Fig. 2. Showing the stratigraphy of the gongola sub-basin 1-mudstone, 2-limestone, 3-sandstone, 4-hiatus, 5-basalt, 6-marine sediments, 7-transitional-marine sediments, 8-continental sediments, 9-basement complex (du-dumbulwa member, df-deba fulani member, gu-gulani member)

2. MATERIALS AND METHODS

Topographic and structural maps of Gombe town and environs were employed in the fieldwork of this research to identify potential areas where Yolde Formation are well exposed. Along these well exposed outcrops identified, lithostratigraphic sections of this Formation outcropping around Gabukka stream (Fig. 3) were systematically logged to record data on lithologic variations, texture, bed geometry, paleocurrents, sedimentary structures and fossil content. Based on facies concept and application of Walters law in conjunction with facies relation

provided by sedimentologic studies on ancient and modern environment, these data were utilized in designating lithofacies assemblages representing particular depositional environment. Paleocurrent measurements were also carried out on the abundant planar and trough crossbedded sandstones and the various orientations determined were used to evaluate provenance and hydrodynamic processes [26]. The dip and strike -plus the azimuth of the crossbeds were measured using compass clinometers in this analysis, and furthermore subjected to tilt correction using the procedure of [26].

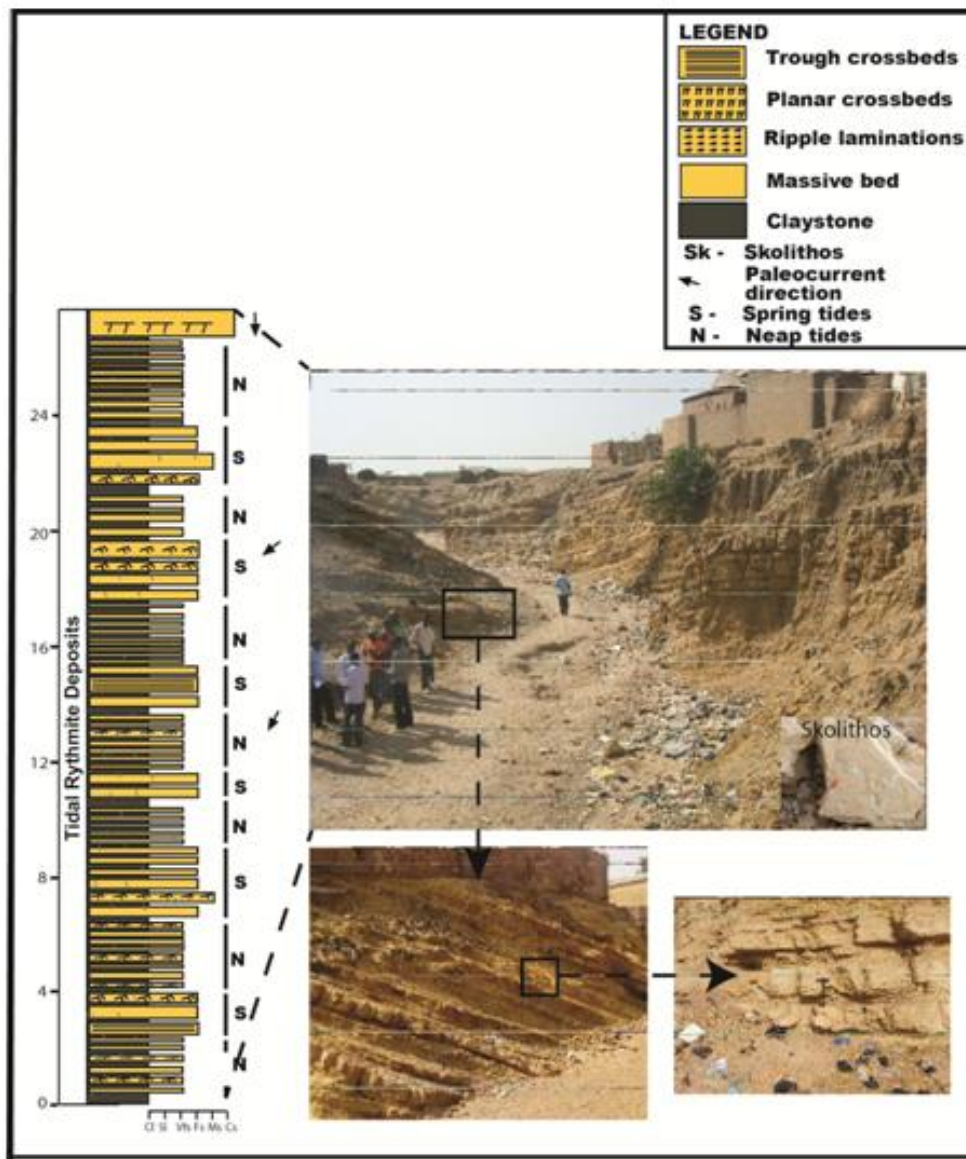


Fig. 3. Showing lithostratigraphic section of the yolde formation at gabukka locality

3. RESULTS

3.1 Sedimentary Facies Characterization

This facies association is formed of successive sequence of horizontally bedded sandstones 2-40 cm interbedded with mudstones of 2-10 cm amalgamating to build succession of over 20 m thick. This package typically displays very fine – fine grained sandstones dominantly of parallel laminated (Sl), ripple laminated (Sr) and massive bedded (Sm) sandstone facies (Fig. 3). Uncommonly, coarser units and planar crossbedded sandstone facies (Sp) are also noticeable often bearing soft sediment deformational structure. The stacked succession of these units are characterized with systematic packages of thickening and thinning clusters of horizontally stratified sandstone units showing marked differentiations in grain sizes, predominantly of fine grained texture in thickening sequences and very fine to siltstones in the thinning packages. Boundaries are generally sharp and bioturbation are scarce, but records of Skolithos ichnofacies are common. Paleocurrent evaluation from the few planar crossbeds and relatively inclined horizontal stratification indicate a southwestern trend.

4. DISCUSSION

These stratigraphic packages displaying sandstones-mudstones symmetry arranged in a relatively predefined harmonic pattern, scarcely bioturbated by Skolithos ichnogenera are indications of a tidal rhythmites deposits [27,28]. These alternations between sand-mud accounts for interchanging dynamics of high and low tide energy and the periodic rarefaction and compression of thicknesses configured to reflect neap-spring tidal sedimentation cycles [29,30]. The mudstone units are product of slack-water periods, appearing as thicker couplets interbedded with thinner plane horizontal sandstone, suggesting deposition during neap tides, commonly occasioned with weak tidal current, thus promoting maximum mud deposition. Contrarily, generation of thicker sandstone beds interbedded with less mudstones packages are skewed to spring tides which is typically accompanied by strong tidal currents [31]. Presences of coarse grained units in these packages are probable indication of attenuated tidal regime deposits, likely driven by superposed wind processes [32], whereas the less common planar crossbedded units are functions of re-circulatory flow from tidal flow

separation associated with large dune migration [33]. The vertical burrows of the low diversity Skolithos ichnofacies consisting of Skolithos is reflective of tidal agitations [34]. Rarity and absence of bioturbations at several intervals are account of high sedimentation rate or changing salinity conditions, most probably induced by high flux of fresh water [35,36].

5. CONCLUSION

The stratigraphic architectural packages of the Yolde formation outcropping at Gabukka locality in the Gongola Sub-basin is composed of five lithofacies that comprises of massive bedded sandstone facies (Sm), planar crossbedded sandstone facies (Sp), ripple laminated sandstone facies (Sr), parallel laminated sandstone facies (Sl) and mudstone facies (Fm). These sandstones facies interbedded with mudstones facies are organized into thickening and thinning packages that reflects spring and neap tides. This implies an intense tide conditions, invariably indicating tide dominated oceanographic regimes, thus the development of the tidal rhythmite stratigraphic architecture in the Cretaceous Yolde Formation at Gabukka locality Gongola Sub-basin of the Northern Benue Trough.

COMPETING INTERESTS

Authors have declared that no competing interests exist.

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